Constraints on the lunar core composition and thermal state from geophysical data and thermodynamic properties of liquid iron alloys

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Constraints on the core of the Moon

- has liquid part (libration from LLR)
 Williams 2001
- cooled fast enough to generate an early magnetic field
- radius 278-440 km and mean density **3900-6750 kg/m³** (seismic data, LLR, tides, induction) ... Garcia 2011; Williams 2014; Khan 2014; Matsumoto 2015, Matsuyama 2016
 - \Rightarrow iron (p_{I-Fe=}7400kg/m³) -rich alloy with a significant amount of light elements
- non candidate light elements: Si (unfavorable redox conditions), O and H (p too low during core-mantle differentiation)
- candidate light elements (formation conditions, bulk composition assumptions, chemical element partitioning between liquid metal and silicates):

S≲**0.5wt%** and **C**≲**5wt%**

... Dasgupta 2009; Chi 2014; Steenstra 2016,2017; Righter 2017

At 5GPa: pFe0.5wt%s~7300kg/m³ and Tliquidus~1950K

 \rightarrow too dense to agree with core density inferences

 $\rho_{Fe5wt\%C} \sim 7100 kg/m^3$ and $T_{liquidus} \sim 1600 K$

 \Rightarrow need larger amounts of light elements

Core modeling

- iron-rich core can have a liquid and solid part
- radius of inner-core radius determined from liquidus temperature and core light element concentration
- need equations of state to compute relevant thermodynamic quantities for modeling interior structure and thermal evolution at relevant pressure and temperature conditions (e.g. density, thermal and chemical expansivity, heat capacity)
- equations of state and liquidus temperature of core alloys should be thermodynamically consistent

Fe liquidus



- liquidus deduced from high pressure data (p>10GPa, Morard et al 2018) is in good agreement with low pressure experimental data (Strong 1973, Liu 1975)
- uncertainty on melting temperature at 1bar ≤10K and at 5GPa ≥150K
- temperature increase along Moon core adiabat *≤* **25K** !

I-Fe equation of state

I-Fe eos are not derived from experimental data acquired at Moon core conditions:

- Anderson et al 1994 (1bar thermoelastic data and 1000GPa shock data)
- Komabayashi 2014 (eos of fcc and hcp Fe and Fe liquidus from Anzellini 2013)
- Dorogokupets et al. 2017 (1bar thermoelastic data, eos of fcc and hcp Fe and Fe liquidus from Anzellini 2013)
- Wagle and Steinle-Neumann 2019 (Ab-Initio)



- along isentropes density differences are below 1% and acoustic velocity differences are below 2%
- but for thermal expansivity, heat capacity, and Grüneisen parameter the differences between the eos' can be quite larger
 ⇒ effect on core temperature
- all isentropes except that of Wagle et al. are less steep than the gradient along the Fe liquidus
 ⇒ bottom-up inner core growing for pure Fe and top-down for Wagle et al.
- all relevant thermodynamic quantities of I-Fe for structure and thermal evolution can be calculated from the eos'

Fe-S core model

Iron-rich liquidus



- highly non-ideal
- can be described with an asymmetric Margules model that has interaction parameters linear in p and T (Buono & Walker 2011)

$$\begin{aligned} G^{l}(X_{\text{FeS}}, p, T) = & (I - X_{\text{FeS}}) \, G^{l}_{\text{Fe}}(p, T) + X_{\text{FeS}} \, G^{l}_{\text{FeS}}(p, T) + \\ & (I - X_{\text{FeS}}) \, R \, T \ln(I - X_{\text{FeS}}) + X_{\text{FeS}} \, R \, T \ln(X_{\text{FeS}}) + \\ & X_{\text{FeS}}(I - X_{\text{FeS}}) \left[X_{\text{FeS}} \, W_{\text{Fe}}(p, T) + (I - X_{\text{FeS}}) \, W_{\text{FeS}}(p, T) \right] \end{aligned}$$

Elastic properties

Morard et al. 2018

Nishida et al. 2016



- densities of liquid Fe-S alloys at 5GPa in agreement with expected average Moon core density (3900-6750 kg/m³)
- Buono & Walker model induces a concentration dependent but (p,T) independent excessive mixing volume that can explain the high pressure density data
- but not the acoustic velocity data

Thermodynamic model

- end-members I-Fe (Komabayashi 2014) and I-FeS
- asymmetric Margules model with pressure dependent excessive volume

$$\begin{aligned} G_{ex}'(X_{\text{FeS}}, p, T) &= X_{\text{FeS}}(I - X_{\text{FeS}}) \left[X_{\text{FeS}} W_{\text{Fe}}(p, T) + (I - X_{\text{FeS}}) W_{\text{FeS}}(p, T) \right] \\ W_{\text{Fe}}(p, T) &= W_{\text{Fe},H} - W_{\text{Fe},S}T + W_{\text{Fe},V} \int_{0}^{p} V_{ex}(p') dp' \\ W_{\text{FeS}}(p, T) &= W_{\text{FeS},H} - W_{\text{FeS},S}T + W_{\text{FeS},V} \int_{0}^{p} V_{ex}(p') dp' \end{aligned}$$

- EoS parameters for FeS (except ρ_{ref} and γ=1.3) (4) and interaction parameters (8) are estimated from liquidus, density, and acoustic velocity data
- ambient pressure density FeS from Kress 2007
- use pseudospinodal eos (Baonza 1995) for V_{ex}
 →V_{ex} decreases with increasing pressure



Data-Fit



Comparison with liquid Fe-10wt%S data used in several Moon core models



- uncertainty on elastic data induces errors on predicted densities of I-Fe-S alloys that are below 5% (taking into account correlations between eos parameters)
- predicted density of Fe-10wt%S is significantly larger than values reported by Sanloup et al. 2000
 ⇒new elastic data and thermodynamic model requires more sulfur to explain average core density

Caveats

- estimated model parameters depend on I-Fe eos
- results depend on reference density of FeS (Kaiura 1979, Kress 2007)
- cannot fit Grüneisen parameter from the data
- cannot assess model predictions for liquidus for compositions above the eutectic because of lack of experimental data
- thermodynamic Fe-S model very much dependent on used elastic data set

Acoustic velocity liquid Fe(Ni)-S



- Jing et al. 2014 Fe(16,30,39)at%S
- Nishida et al. 2016 Fe(20,43,50)at%S
- Terasaki et al. 201x
 Fe10at%Ni(17,30)at%S
- weak dependence on temperature (curves are on isotherms 1900K)

 \Rightarrow Inconsistency between different studies

Effect of elastic data on thermodynamic model (5GPa, 2000K)



- density of liquid solution based on Morard 2018 density data depends weakly on acoustic velocity data
- but predicted acoustic velocities are quite different ...

Preliminary Fe-C core model



- at ~5GPa C saturation in I-Fe ≲7wt%C
- assume ideal mixture of I-Fe (Komabayashi 2014) and I-Fe3.5wt%C (Shimoyama 2016)

 \Rightarrow Density of liquid Fe-C significantly larger than 3900-6750 kg/m³



Moon models

- for illustration use the mantle density model of Weber et al. (2011)
- agree with the latest estimate of average shell moment of inertia (MOIs = 0.393112 ± 0.000012, Williams et al., 2014)
- core thermal evolution model based on Davies et al. 2015 and mantle evolution model based on Morschhauser et al. 2011

Structure functions: Fe-S and Fe-C

Core-mantle boundary Temperature: 1920K



- the required amount of S to match the expected core density of the Moon is significantly larger than what is deduced form core formation models
- the weight fraction of C in liquid Fe-C is below 7wt% and the density of such an alloy is significantly above what is expected for the core of the Moon

Structure functions Fe-S (Morard-Nishida)



- models with the Weber et al. mantle cannot have an inner core (at 1σ)
- inner core possible if $r_{cmb} \approx 295$ km and $x_S \approx 5$ wt%
- bottom-up inner core growing requires x_S≤4.2wt%
- to avoid present-day lower mantle melting T_{cmb}≲1920K (Hirschmann et al. 2012)

Thermal evolution with bottom-up inner core formation



- all required thermodynamic quantities for core thermal evolution are computed from core model of this study (density, heat capacity, latent heat of crystallization, thermal- and chemical expansivity)
- main power and entropy source is latent heat
- timing of dynamo agrees with expected period of occurrence but predicted surface magnetic field is significantly below the 20-110µT that are expected to explain the lunar magnetic records (Tikoo et al 2017)

Conclusions

- melting data and new elastic data about Fe-S alloys can be described with a non-ideal mixing model that has a pressure dependent excess volume
- Fe-S models with an inner core and without a whole snowing liquid core cannot be much colder than ~1825K and those models have less than ~4.3wt% of sulfur
- models with an inner-core agree with the timing of occurrence of the lunar dynamo but not with the expected field strength
- but the amount of S required to match the expected core density is too large to allow for bottom-up inner core formation and significantly above what is expected from core formation models
- C cannot be the only light element in the core because even at saturation concentration such a Fe-C alloy is denser than the expected Moon core density

Thermodynamic quantities



Thermal evolution





Core composition

• Steenstra 2016

depletion of siderophiles elements in the core can be explained if S>8wt% but S abundance in mantle implies S poor core but V and Cr abundances require either a S poor core with a differentiation temperature >3100K but core mass~2.5wt% (my models<1.3wt%) or S~8.5wt% with Tdiff~2200K difficult to explain anything with mCore<2.3wt%

• Steenstra 2017

to explain observed S,Se,Te in silicate Moon requires fully molten Moon at core-mantle equilibration

• Steenstra 2017

metal-silicate partitioning of C,S,Ni and BSM (are they reliable?) abundances of S and C S<0.16wt%, main light element carbon up to \sim 4.8wt%

• Righter 2017

Structure functions: Fe-S and Fe-C

Core-mantle boundary Temperature: 1920K



- models based on elastic data of Terasaki et al 2019 require less sulfur at a given core radius to explain the density of the core
- C weight fraction in liquid Fe-C is below 7wt% and
- core densities of Fe-C models are significantly larger than the expected core density of the Moon